

# Fourth Annual Conference on Carbon Capture & Sequestration

*Developing Potential Paths Forward Based on the  
Knowledge, Science and Experience to Date*

*Ocean Sequestration*

## Progress in Carrying Out Ocean CO<sub>2</sub> Perturbation Experiments

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May 2-5, 2005, Hilton Alexandria Mark Center, Alexandria Virginia

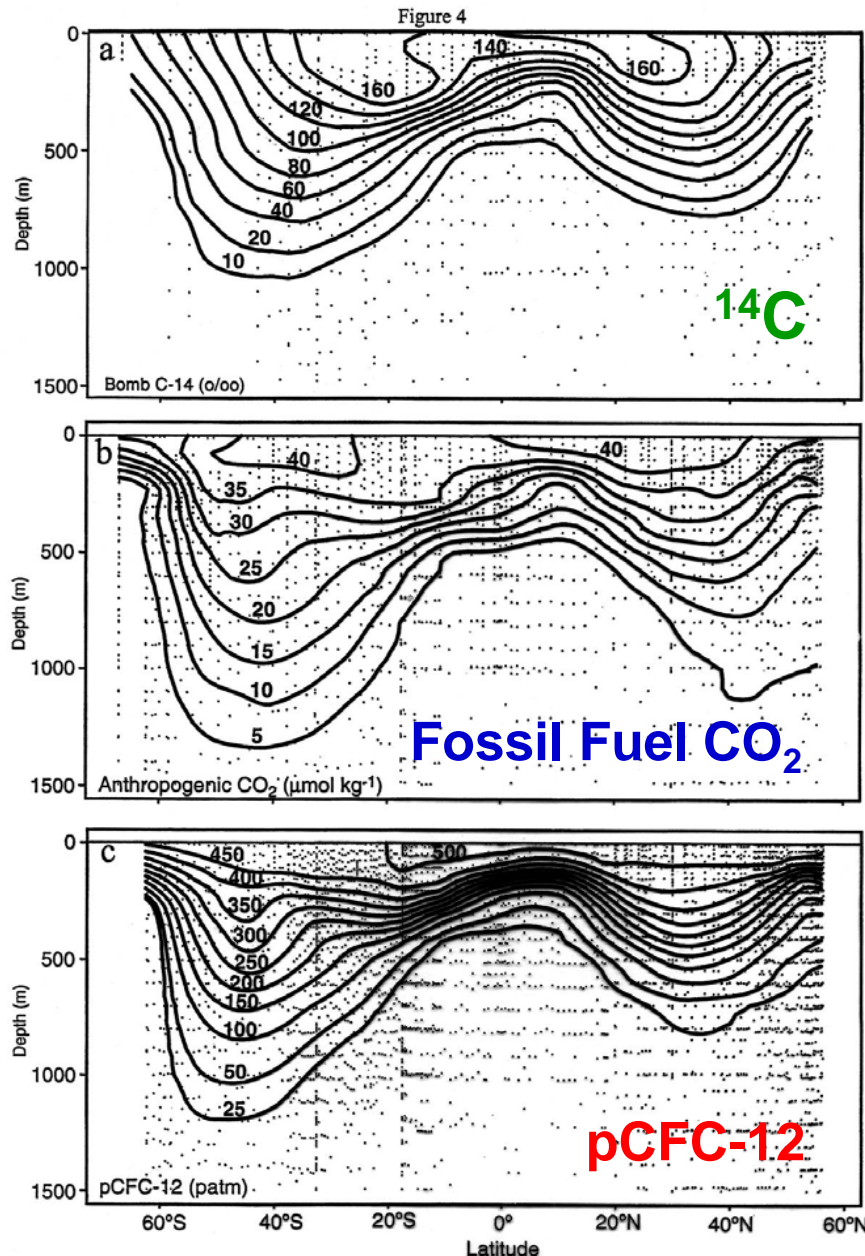


# Ocean Carbon Sequestration

- First proposed by Marchetti, 1977:
  - Ocean is a big place =  $1.4 \times 10^9 \text{ km}^3$ .
  - Already contains 50 times atmospheric burden of  $\text{CO}_2$ .
  - Seawater is strongly buffered and is slightly basic so  $\text{CO}_2$  uptake is the “natural” reaction.
- Current situation:
  - 1/3rd of anthropogenic emissions (2 Gt C equivalent to 7 Gt  $\text{CO}_2$ ) is absorbed annually.
  - As atmospheric levels rise absorption will increase.
  - ~ 80-90% of anthropogenic emissions will go there.
  - Over 500 Gt  $\text{CO}_2$  absorbed to date.



# Ocean CO<sub>2</sub> “Disposal” Today



See: Sabine et al. (2004) in *Science*.

JGOFS/WOCE survey data. Pacific meridional section.

Fossil fuel signal has penetrated to >1000m. Surface values approach 50 μmol/kg (2.2 mg/kg).

The global ocean inventory of anthropogenic CO<sub>2</sub> from 1800 to 1999 is:

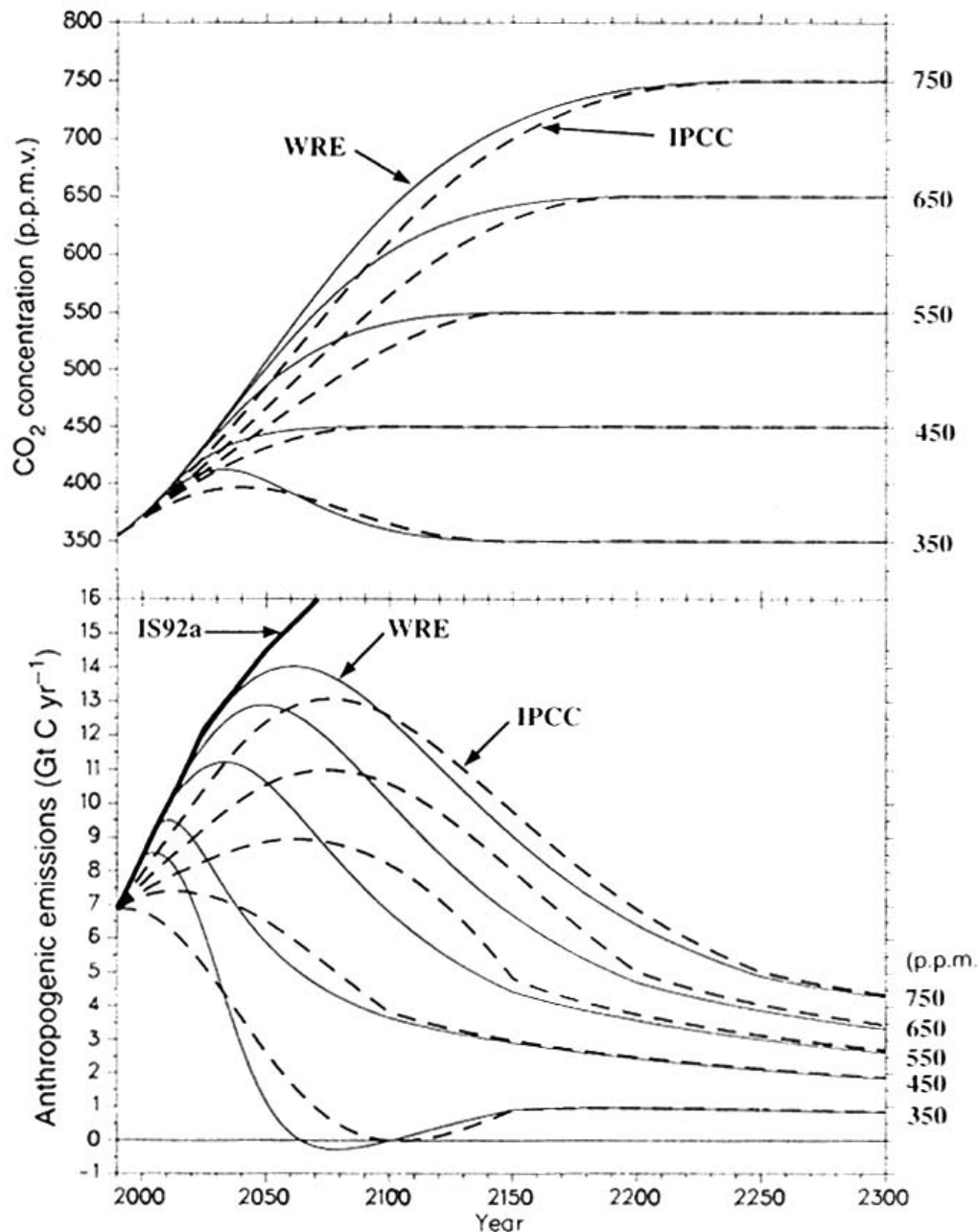
$$155 \text{ Gt C} = 568 \text{ Gt CO}_2.$$

Global surface ocean CO<sub>2</sub> disposal is now about 20-25 million tons per day.

# Ocean Carbon Sequestration

- Problems:
  - Natural carbon sequestration processes are slow.
  - If uptake faster, atmospheric CO<sub>2</sub> rise minimized.
  - Internal circulation of ocean is slow and stratified.
  - 1/2 of anthropogenic CO<sub>2</sub> uptake has remained in the upper 400m of the ocean.
  - Sea surface pH already reduced by 0.1
- Technological fixes:
  - Iron fertilization to enhance biological pump.
  - Purposeful sequestration by industrial means.





• *From Wigley et al. (1996)*

• *Emission trajectories required to achieve stabilization of atmospheric CO<sub>2</sub> levels at various values incorporating “best” economic choices.*

• *In order to achieve stabilization at 550 ppmv departures from current trends of about 3.67 billion tons of CO<sub>2</sub> per year are required by 2025, and about 14.7 billion tons per year by 2050.*

• *This may be achieved by conservation, substitution, or sequestration strategies.*

## The evolving chemistry of surface sea water under “Business as Usual”

Time	pCO <sub>2</sub>	Total CO <sub>2</sub>	pH	HCO <sub>3</sub> <sup>-</sup>	CO <sub>3</sub> <sup>2-</sup>	H <sub>2</sub> CO <sub>3</sub>
yr.	μatm	μmol kg <sup>-1</sup>		μmol kg <sup>-1</sup>	μmol kg <sup>-1</sup>	μmol kg <sup>-1</sup>
1800	280	2017	8.191	1789	217	10.5
1996	360	2067	8.101	1869	184	13.5
2020	440	2105	8.028	1928	161	16.5
2040	510	2131	7.972	1968	144	19.1
2060	600	2158	7.911	2008	128	22.5
2080	700	2182	7.851	2043	113	26.2
2100	850	2212	7.775	2083	97	31.8

*Under IPCC “Business as Usual” the pH of surface sea water drops by 0.4 pH units by 2100. CO<sub>3</sub><sup>=</sup> in surface water drops by 55% from pre-1800 values. It will be hard to meet even these goals.*





# MBARI Small-scale CO<sub>2</sub> experiments

- We have been doing small-scale CO<sub>2</sub> release experiments since 1996:
  - Kinetics of CO<sub>2</sub> & CH<sub>4</sub> hydrate formation.
  - Rates of liquid CO<sub>2</sub> & CO<sub>2</sub> hydrate dissolution.
  - Amount and extent of pH disturbances.
  - In collaboration with biologists (J. Barry and M. Tamburri), have looked at the impact on biota.
- Problems:
  - Limited control of experimental conditions.
  - Observed novel but not reproducible results.



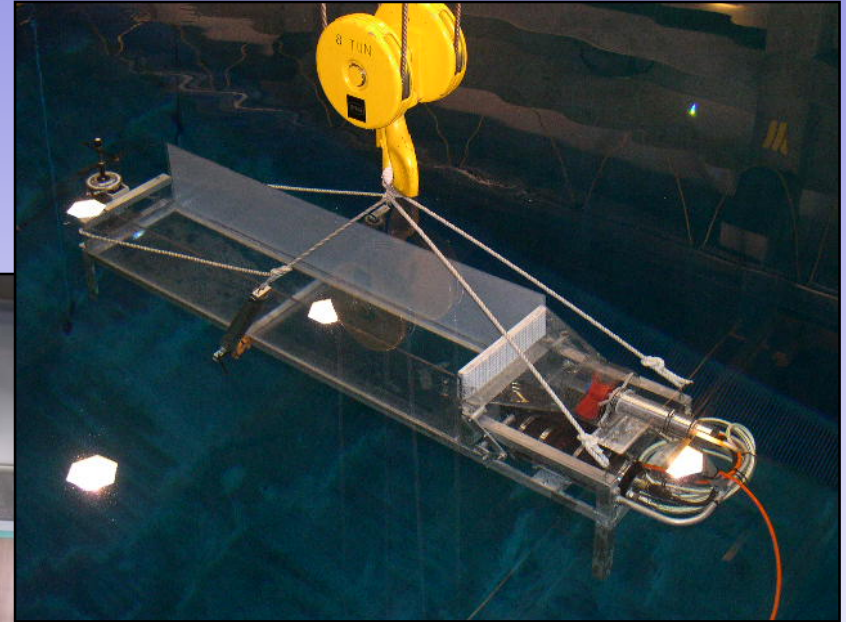
# OACE Research Interests

- MBARI:
  - Develop and deploy new technology for controlled seafloor CO<sub>2</sub> release experiments.
  - Liquid CO<sub>2</sub> and CO<sub>2</sub> hydrate dissolution rates.
  - CO<sub>2</sub> hydration / dehydration kinetics.
- University of Bergen:
  - Fluid-dynamics & chemistry of liquid CO<sub>2</sub> in deep sea.
- NMRI:
  - Develop and conduct pressure tank experiments.
  - Hydrate skin strength and growth rates.
  - Liquid CO<sub>2</sub> / sediment interactions.

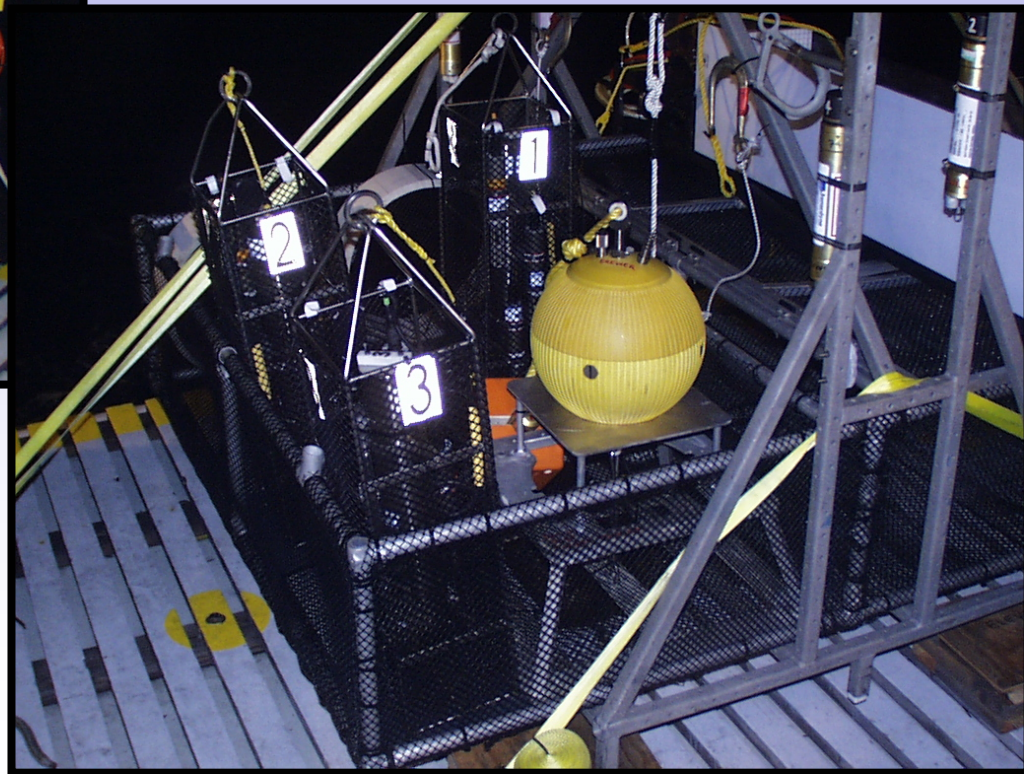
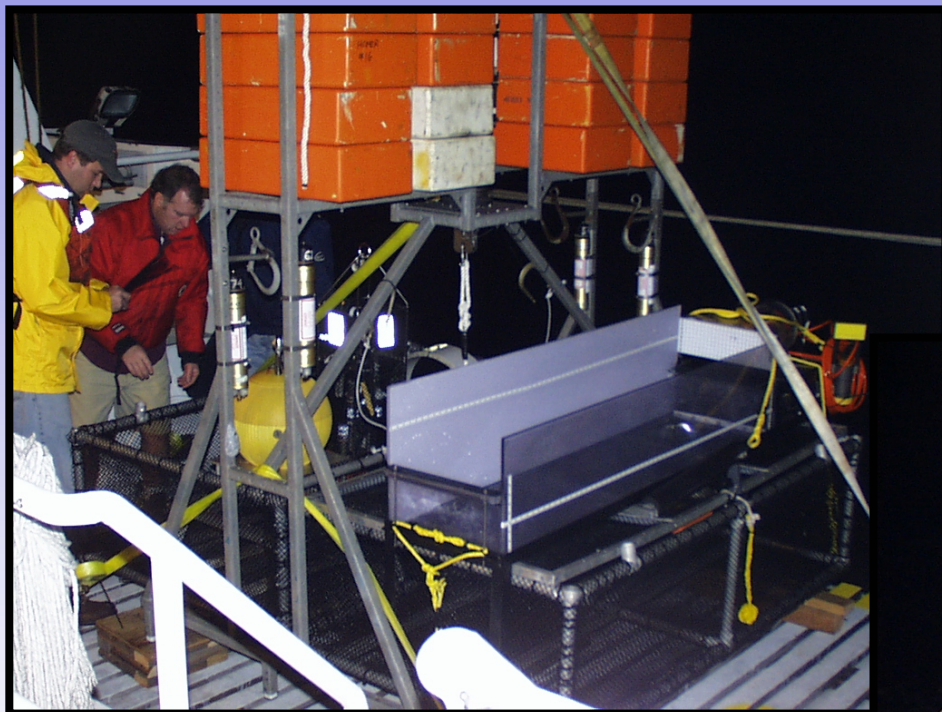




# Benthic Flume



# Equipment elevator



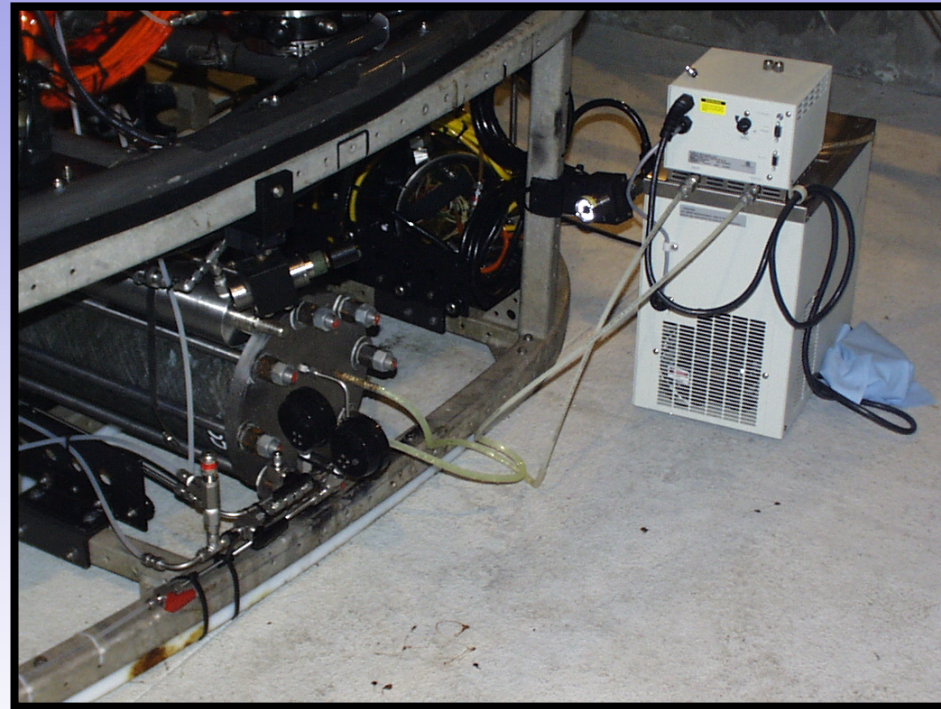
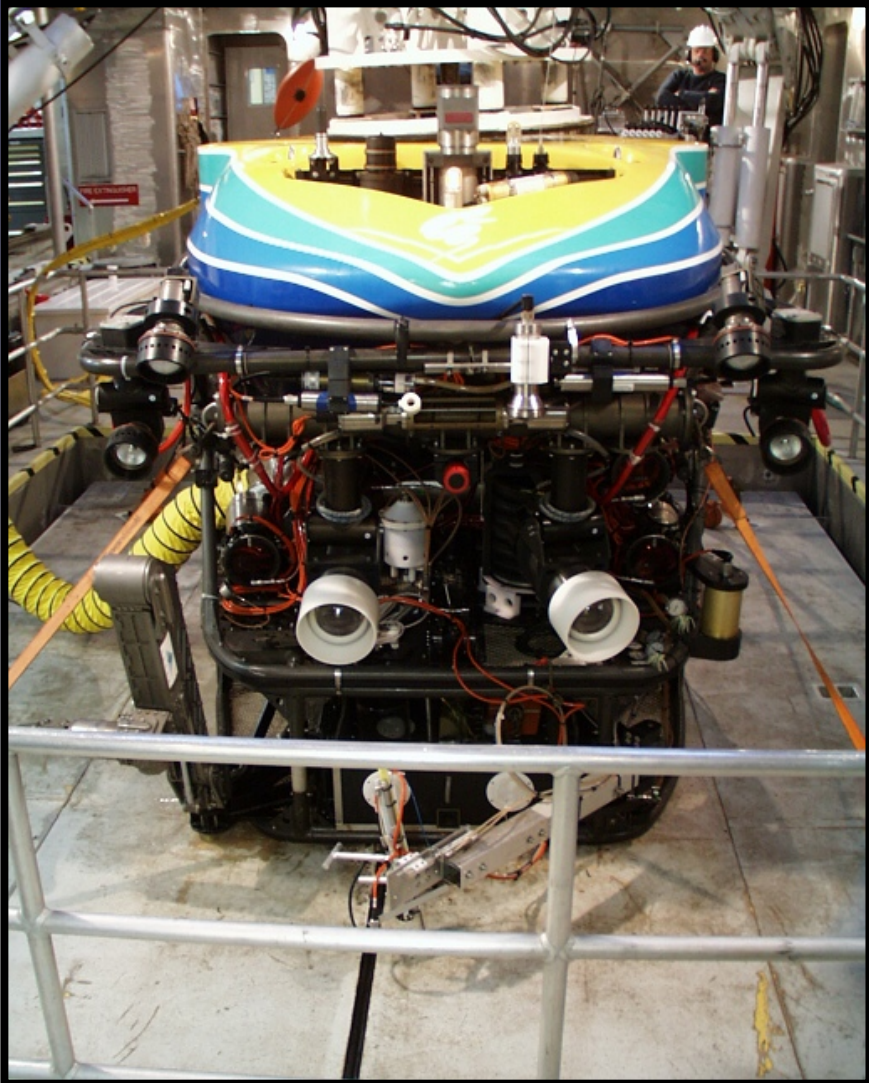


# Elevator deployment



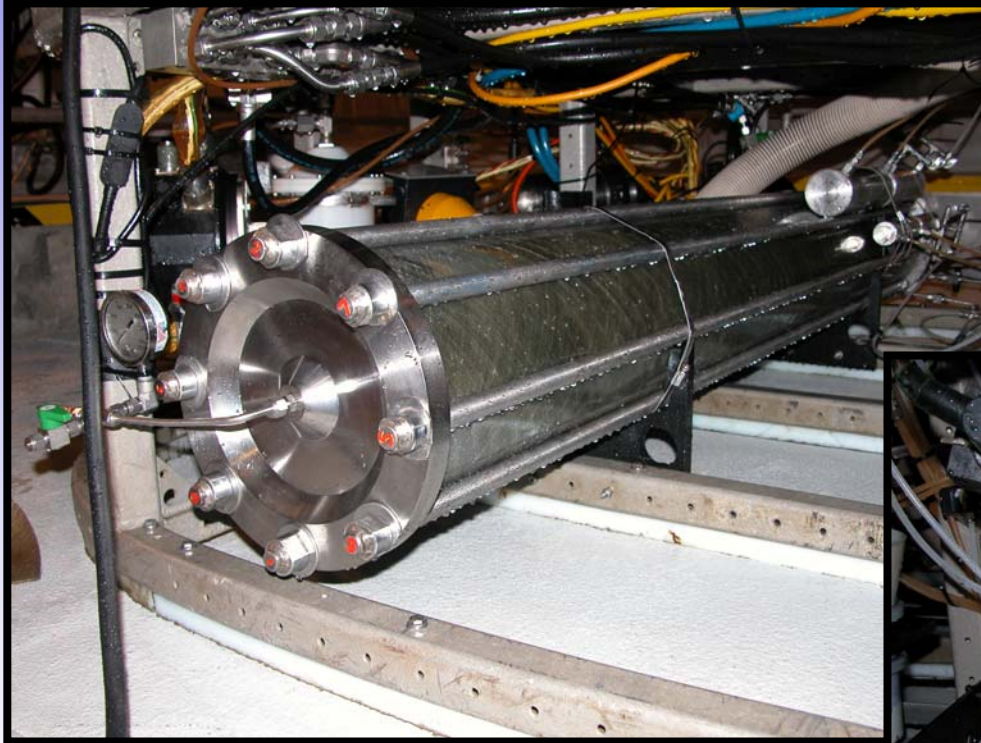


# ROV Tiburon



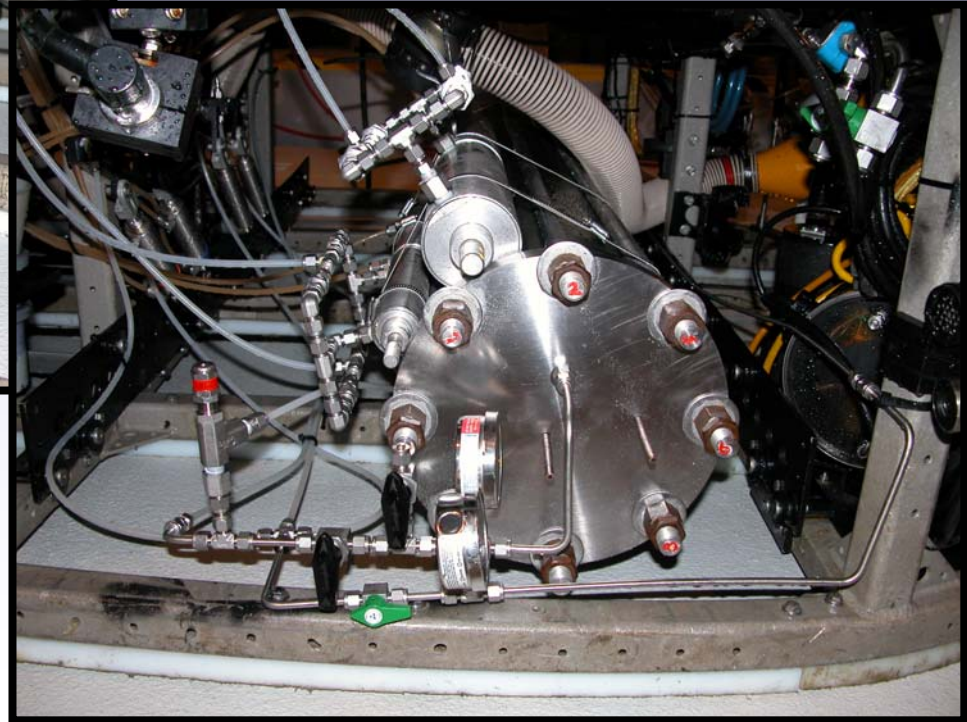


# 56L CO<sub>2</sub> Accumulator



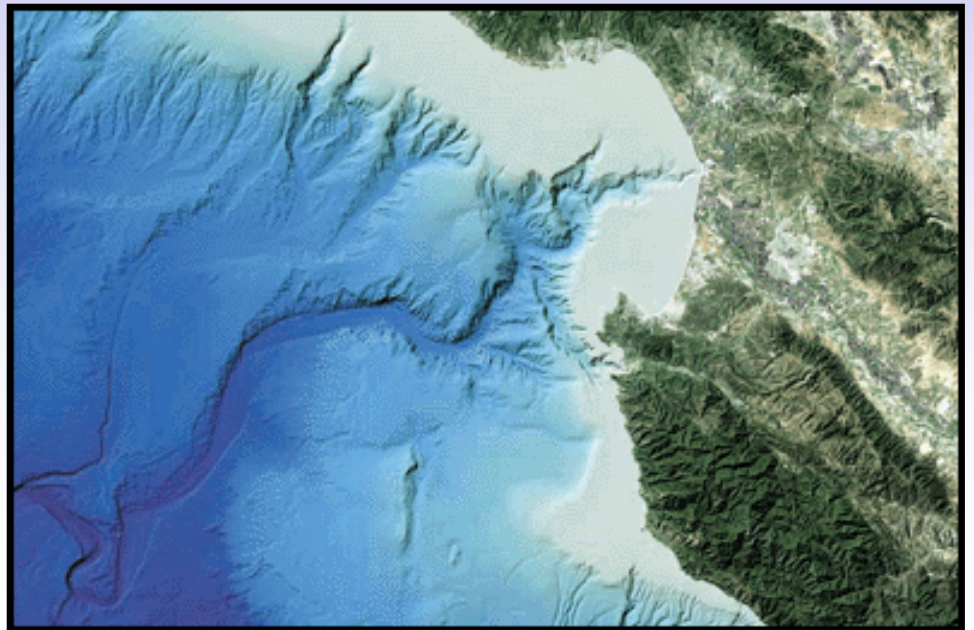
•SW side

•CO<sub>2</sub> side





# Experiment conducted in Monterey Bay



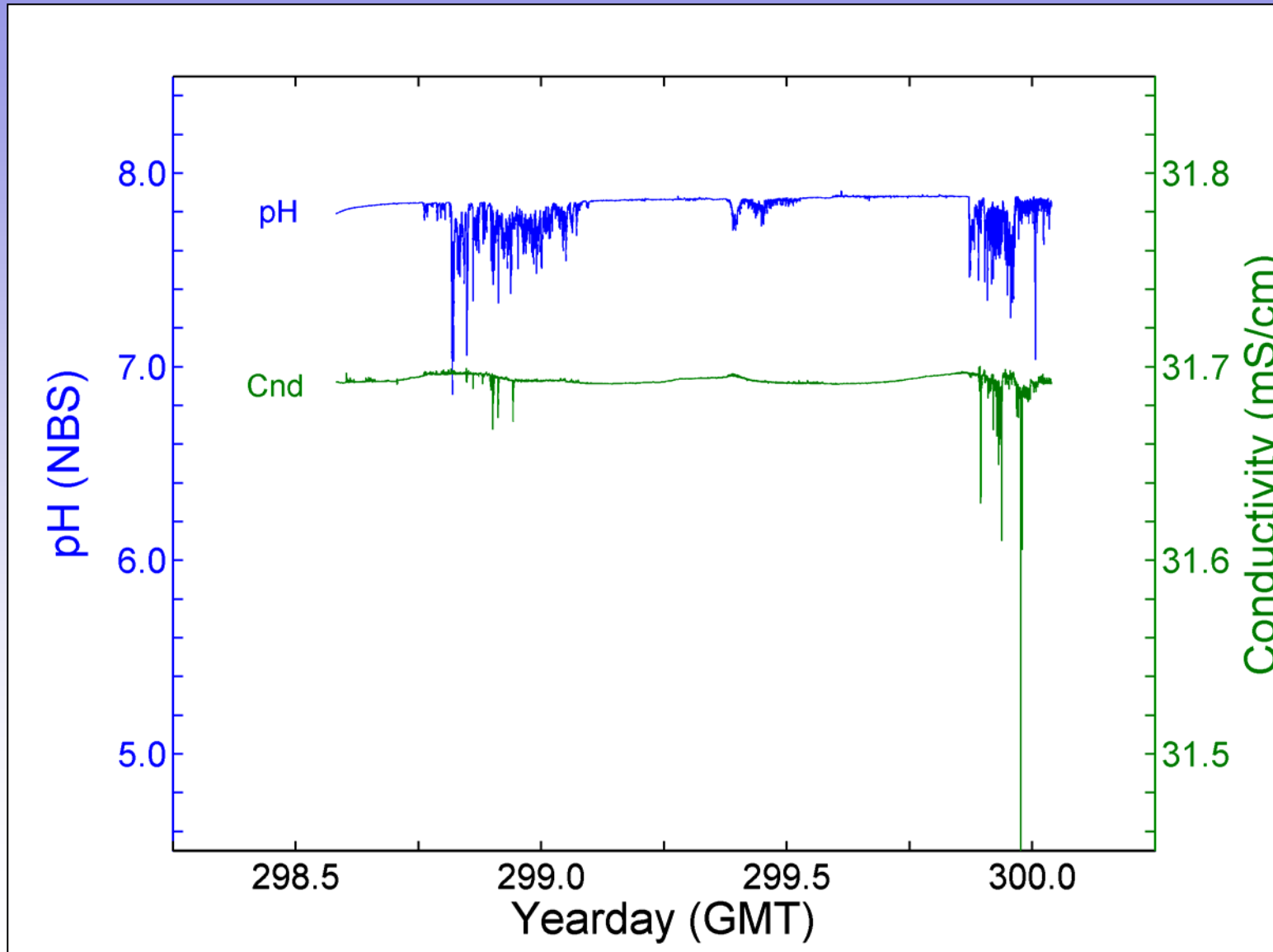
# Video Highlights

- Dye injection experiments at various flowrates.
  - pH indicating dye = phenol red.
  - yellow (6.6)  $\leftrightarrow$  red / purple (8.0).
- Wave generation by “bottom currents.”
  - Slow flow  $\rightarrow$  ripples.
  - Faster flow  $\rightarrow$  breaking waves and droplets:  
“CO<sub>2</sub> spray”

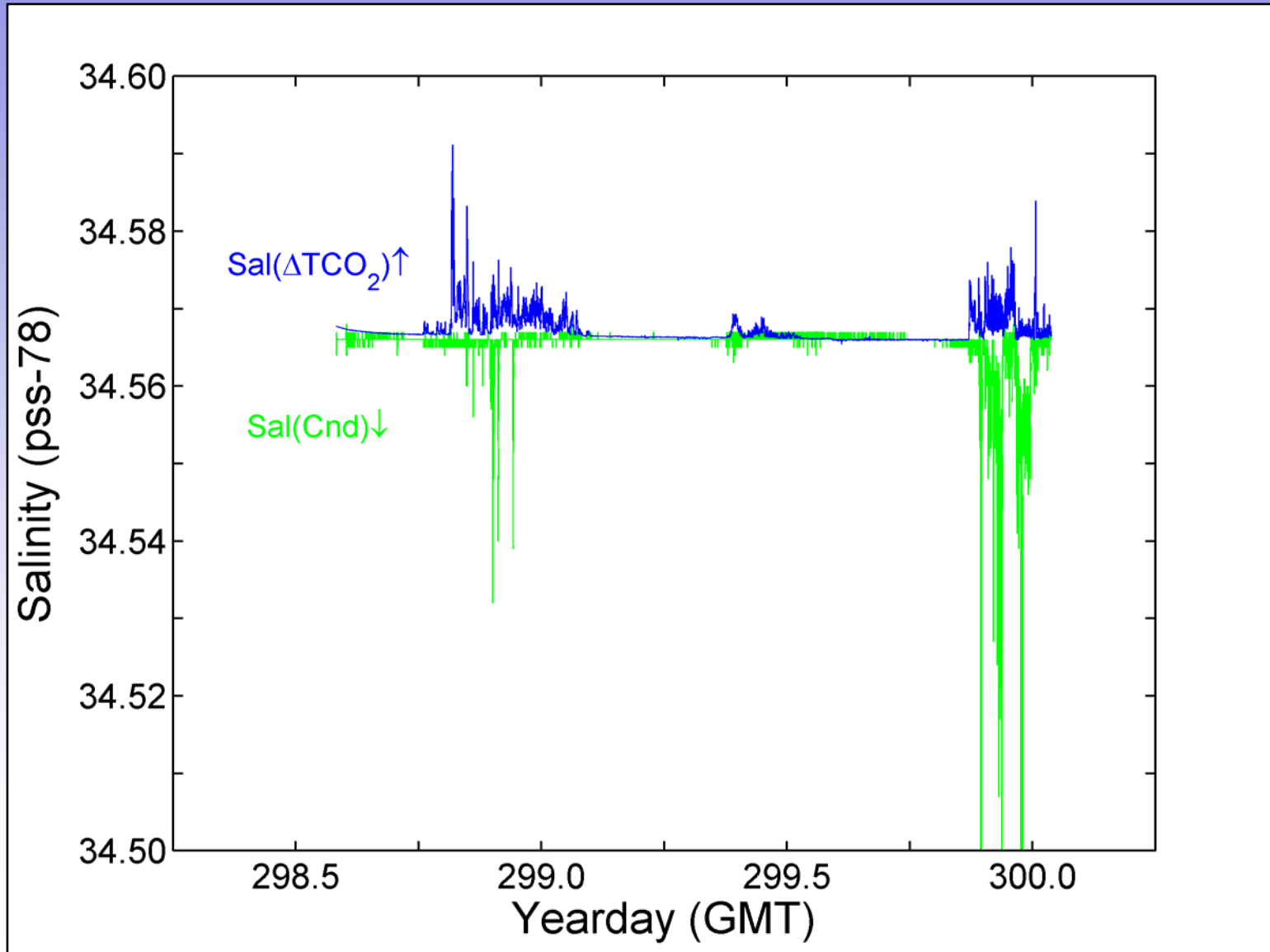




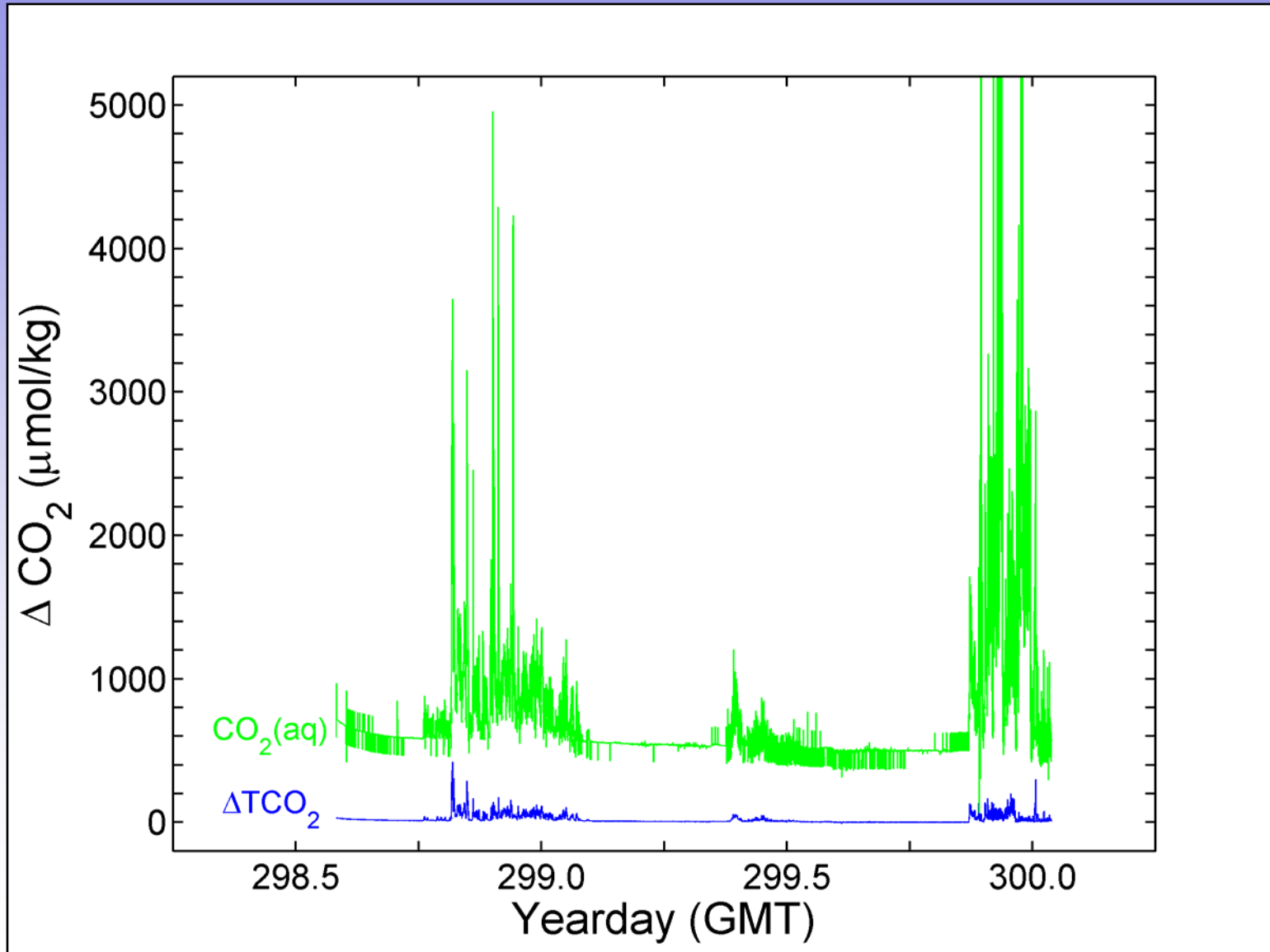
# CTD data record downstream from flume.



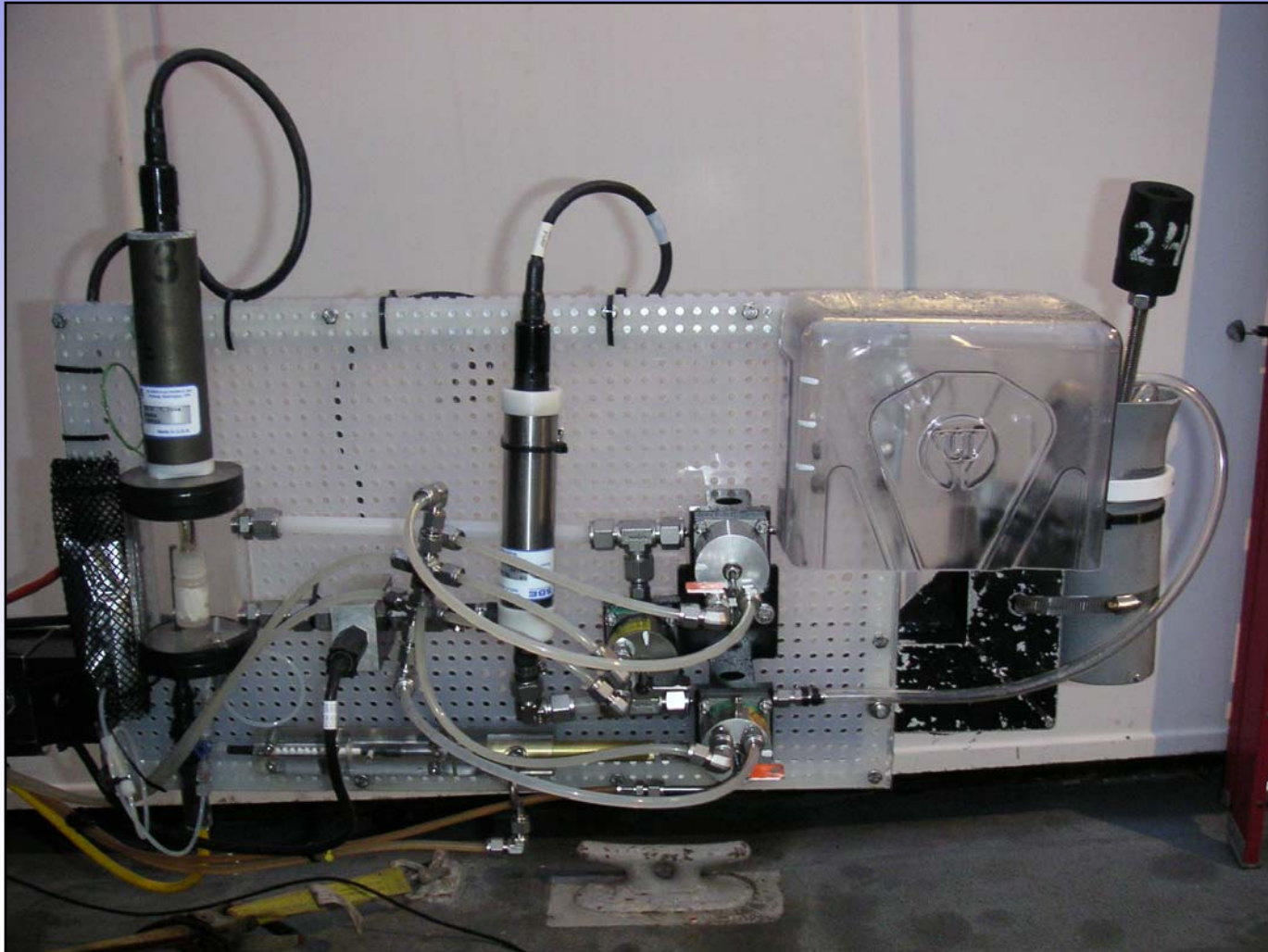
# Salinity observed and calculated downstream from flume.



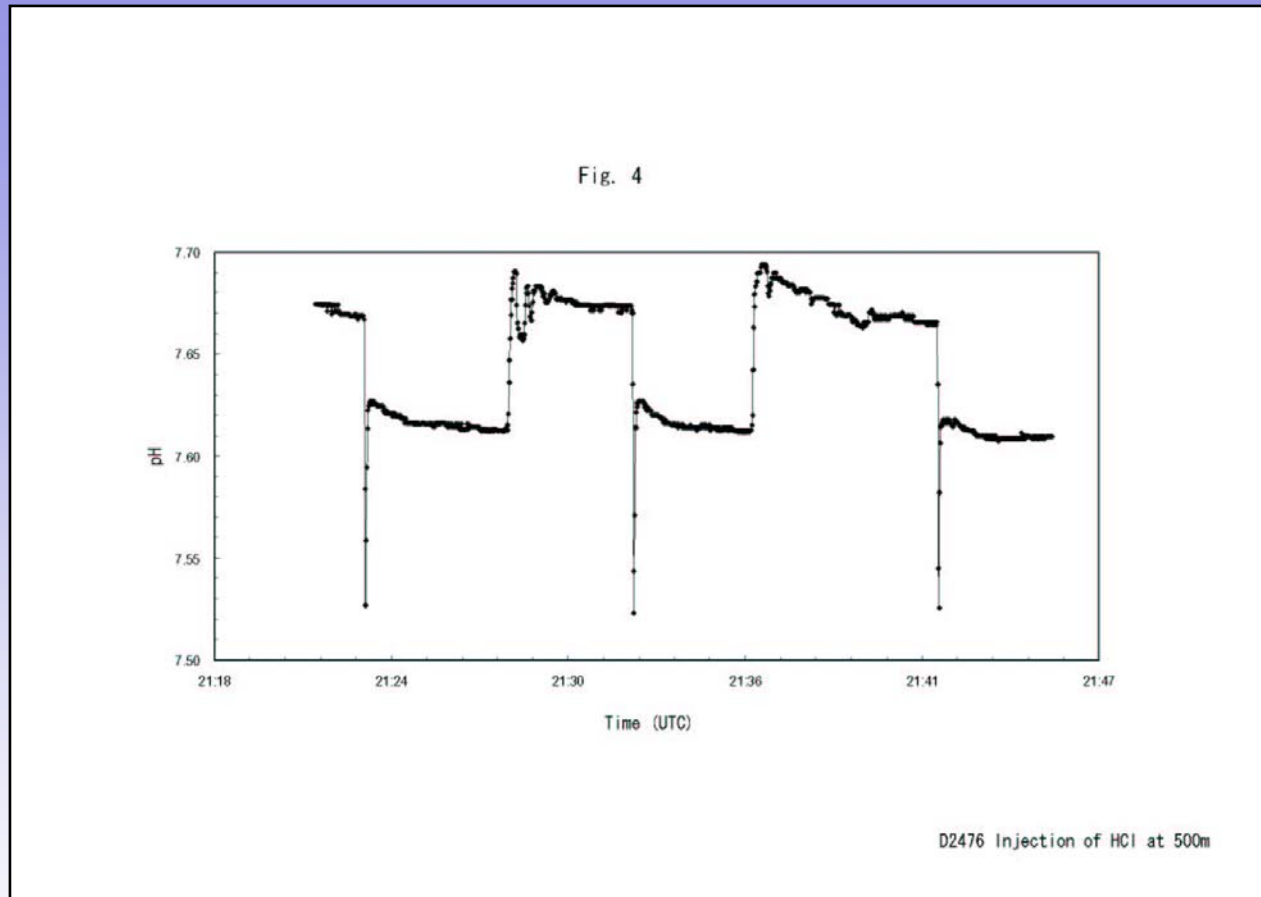
$\Delta\text{CO}_2$  concentration downstream from flume.



# Seawater re-circulation cell

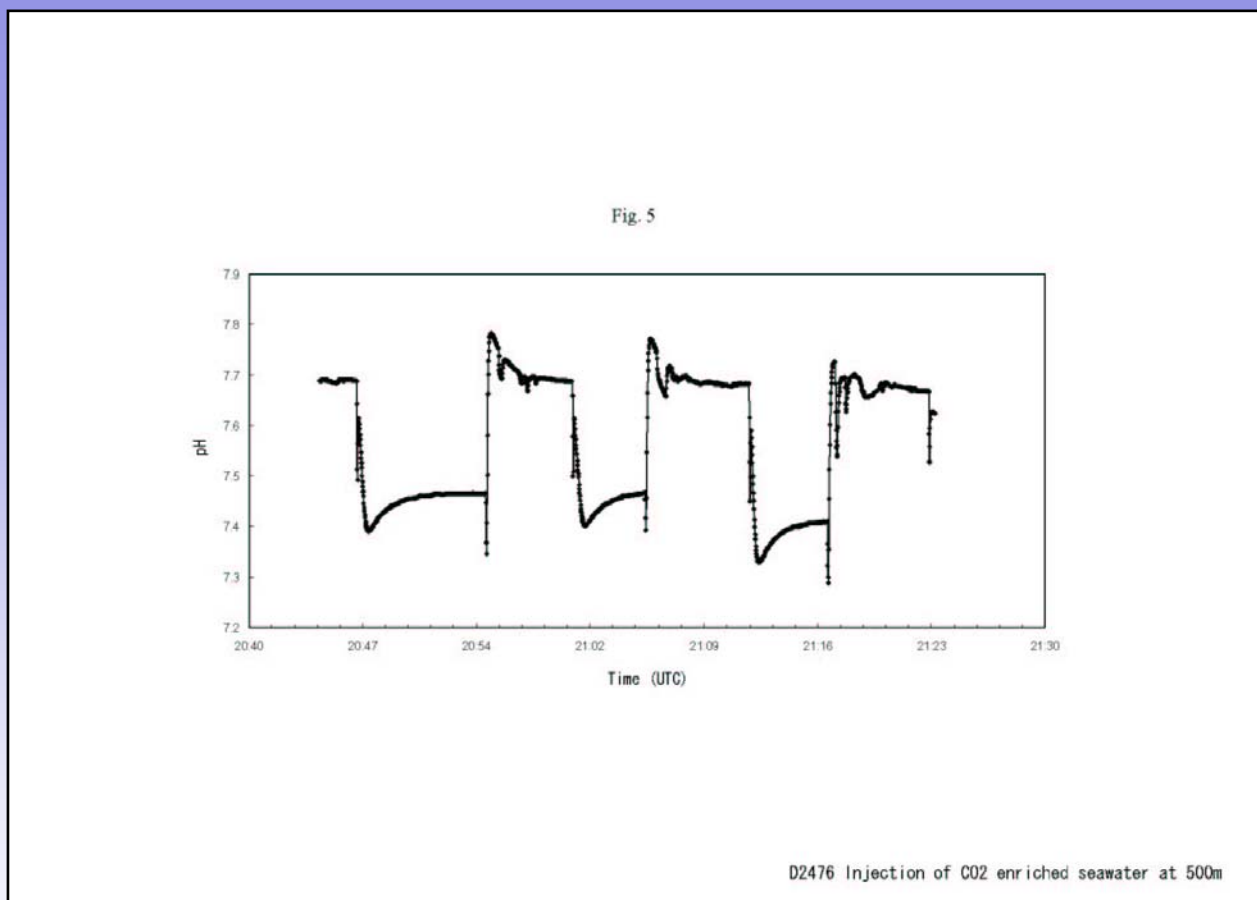


***Results from a 500m depth acid injection experiment to test the feasibility of in situ controlled pH perturbation experiments.***



***The data are consistent with formation of an intense cloud of  $\text{CO}_2$  at the injection tip, followed by rapid mixing and hydration of the excess  $\text{CO}_2$  to form  $\text{HCO}_3^-$ . This reaction has a volume decrease, and is accelerated by pressure.***

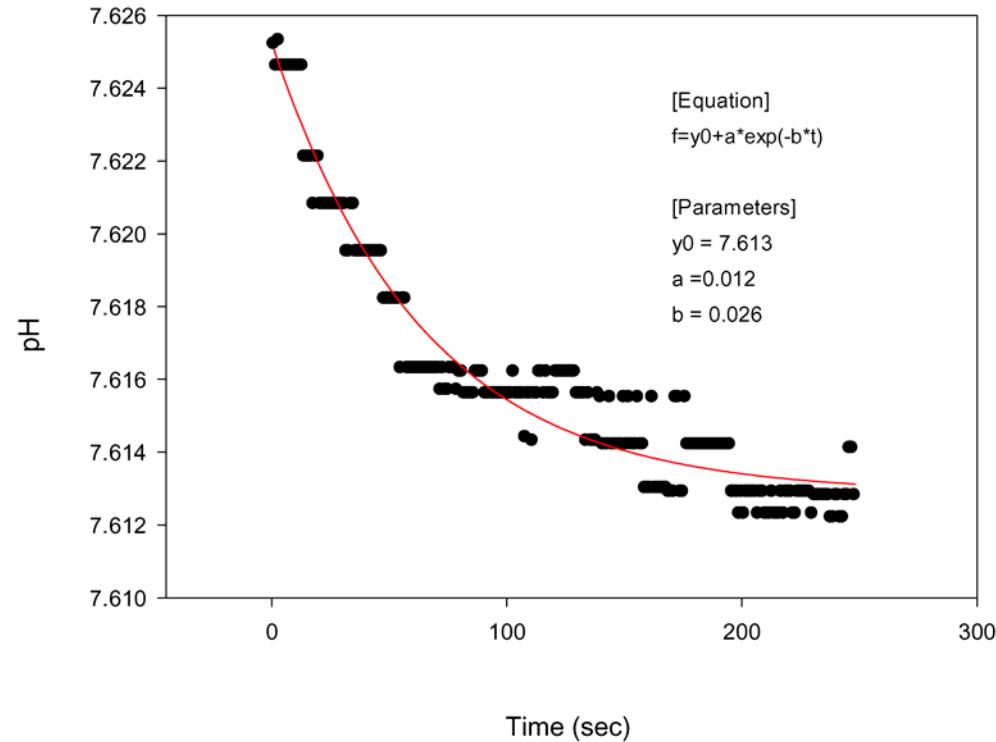
***Results from an injection of  $\text{HCO}_3^-$  rich water in contact with a  $\text{CO}_2$  pool.***



***The shape of the curve differs from acid injection and is consistent with removal of excess  $\text{HCO}_3^-$ . This requires a volume increase, and the reaction at depth is slower than for acid addition – but still far faster than at 1 atm.***

**pH versus time curves for an experiment at 500m depth (6°C) for both acid addition (upper left) and  $\text{HCO}_3^-$  rich water addition (lower right).**

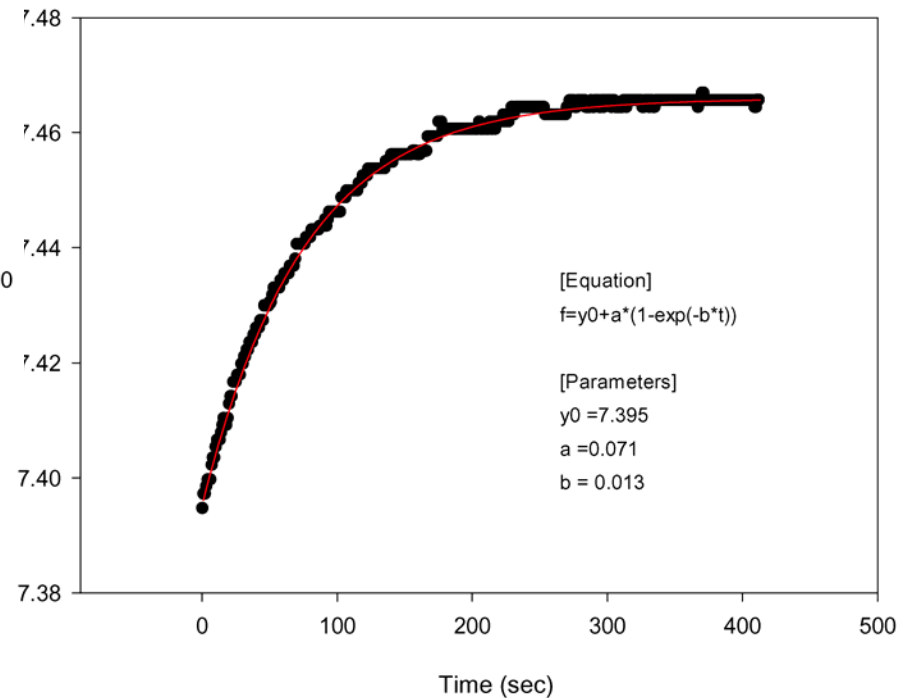
Fig.A



**The initial rapid change is not fit, only the slower exponential decay at the end.**

**Note the reversed shape, and the slower response of the  $\text{HCO}_3^-$  curve.**

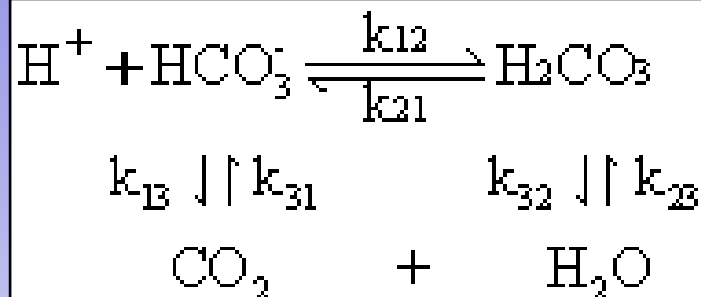
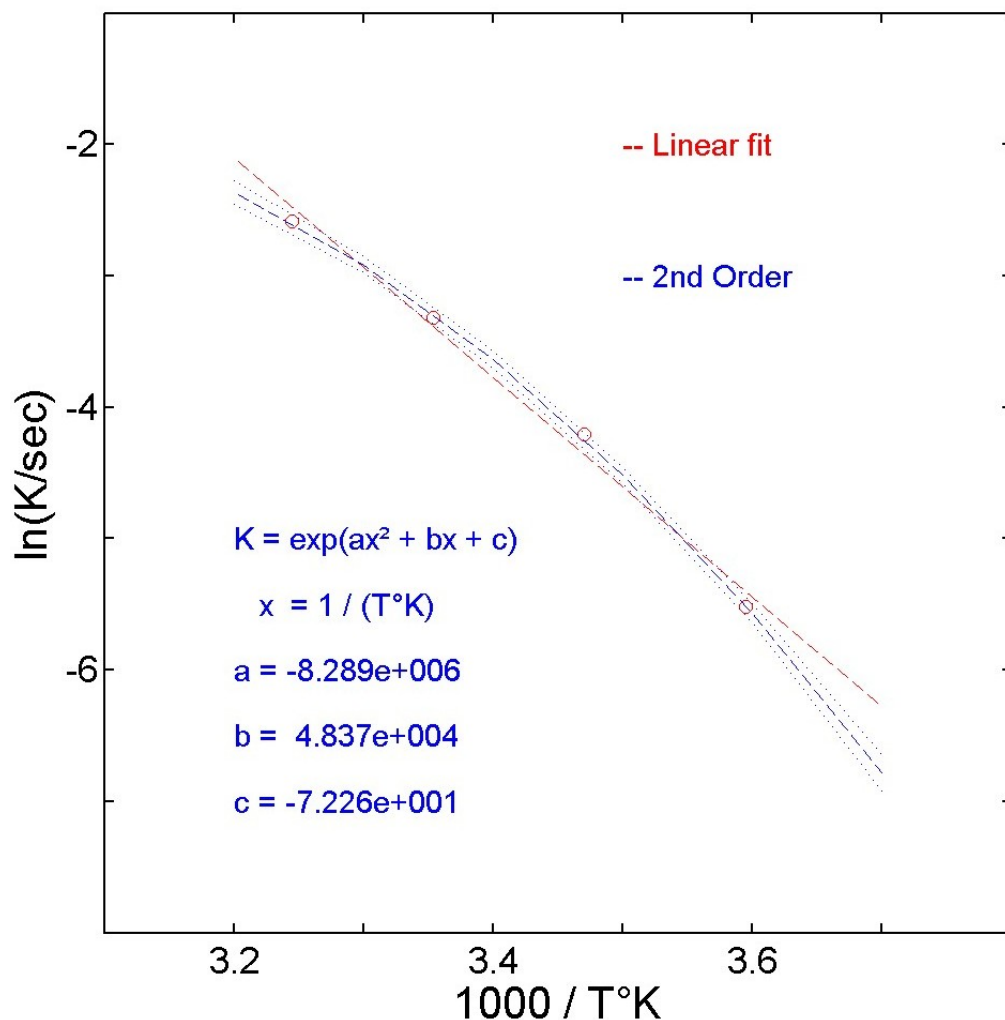
Fig. B





# The CO<sub>2</sub> hydration rate constant (Johnson, 1982)

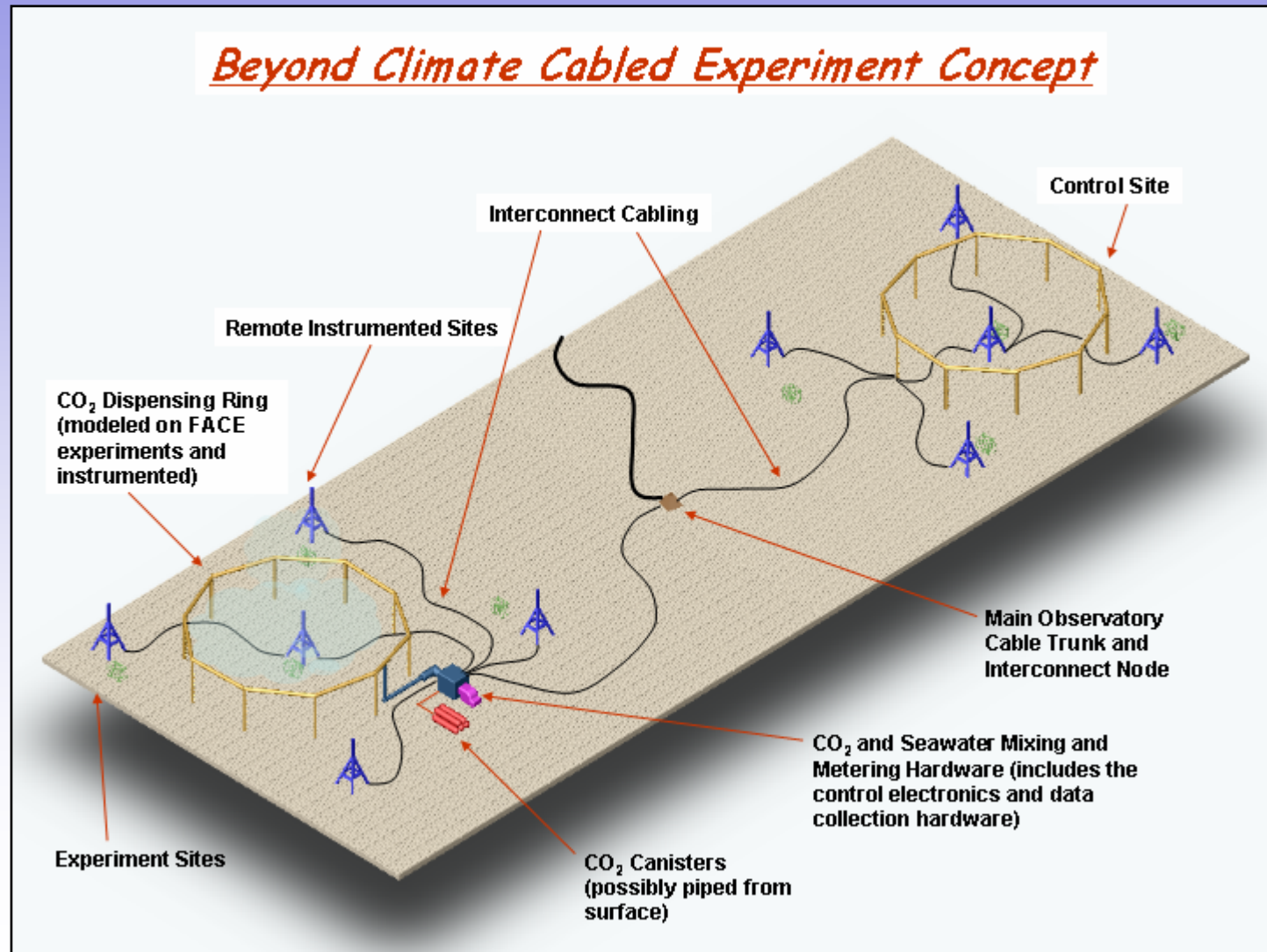
CO<sub>2</sub> Hydration Rate



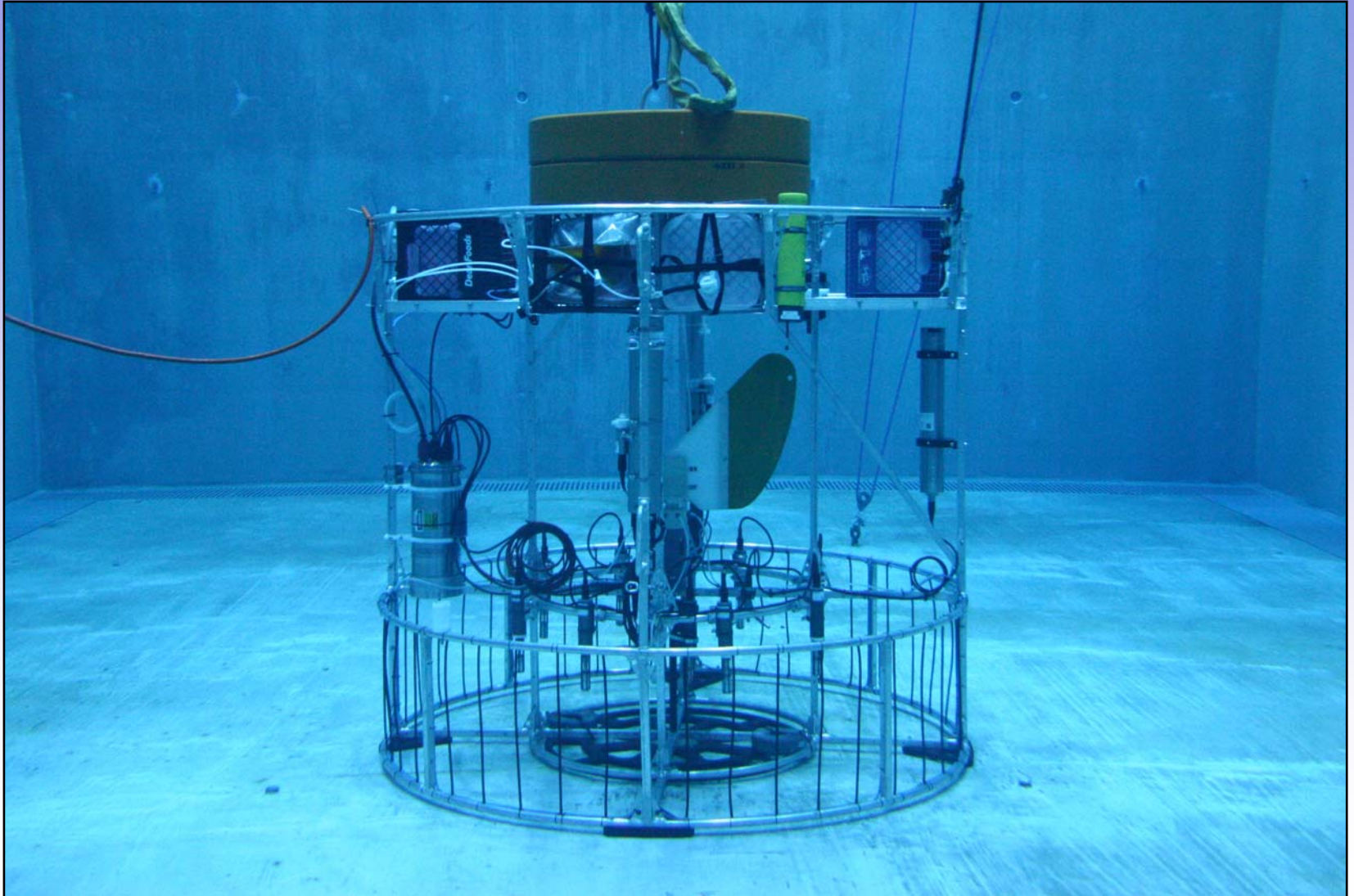
***These data are for 1 atmos.  
They imply a time to equilibrium  
of ~32 minutes at 1.6°C, and 17  
minutes at 6°C; this rxtn time is  
far beyond our observations.***

***However the dissolution of CO<sub>2</sub>  
produces a strong -Δ V, and  
thus the effect of pressure is to  
shift the equilibrium to the  
hydrated state – but the rates  
at pressure are unknown.***

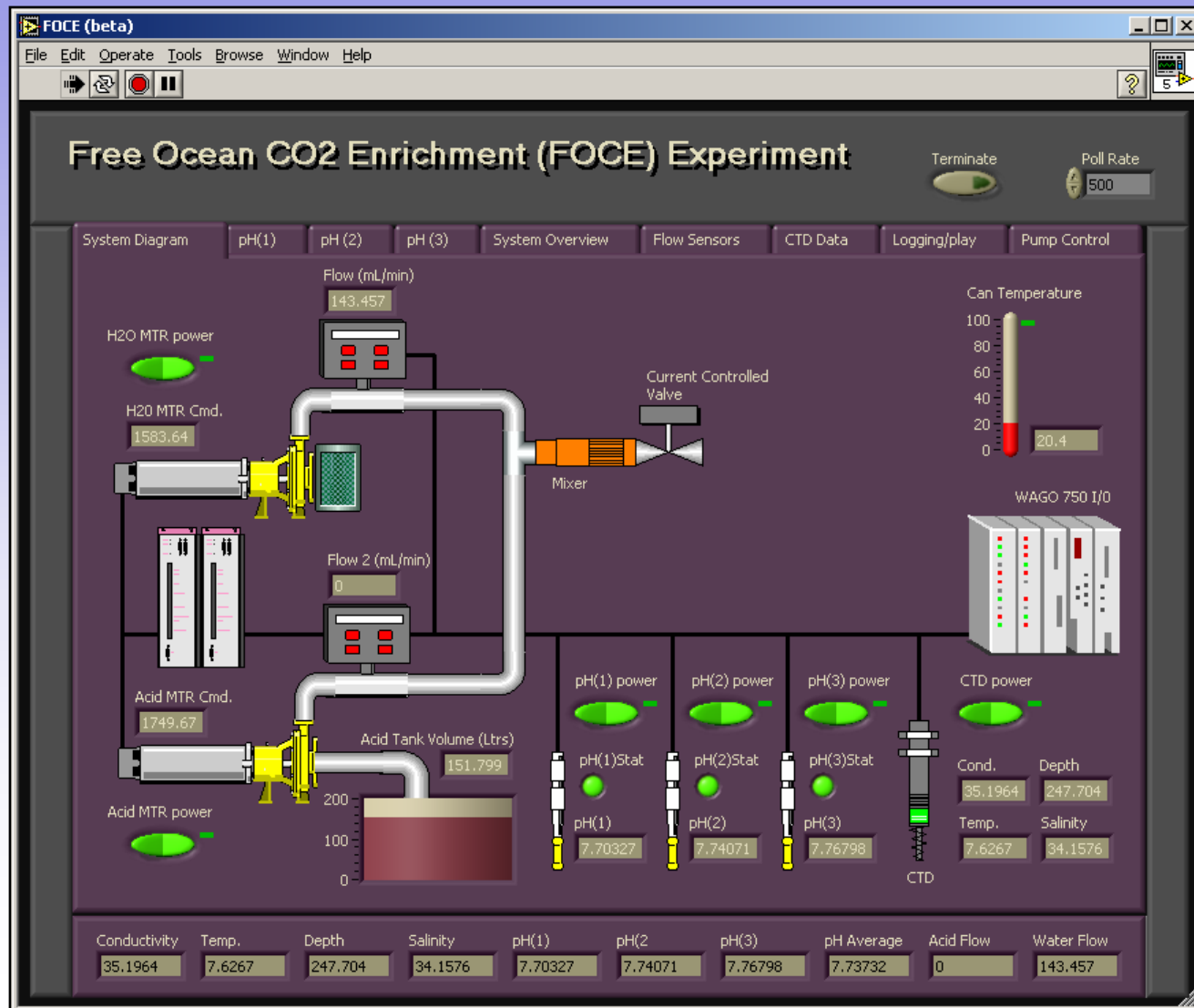
*If we are to understand the science of a lower pH ocean we have to carry out predictive experiments. A concept sketch is shown here with supply of either acid or CO<sub>2</sub> to a set of experimental sites in much the same way that experiments are carried out on land. There are fundamental challenges in this.*



# FOCE Prototype

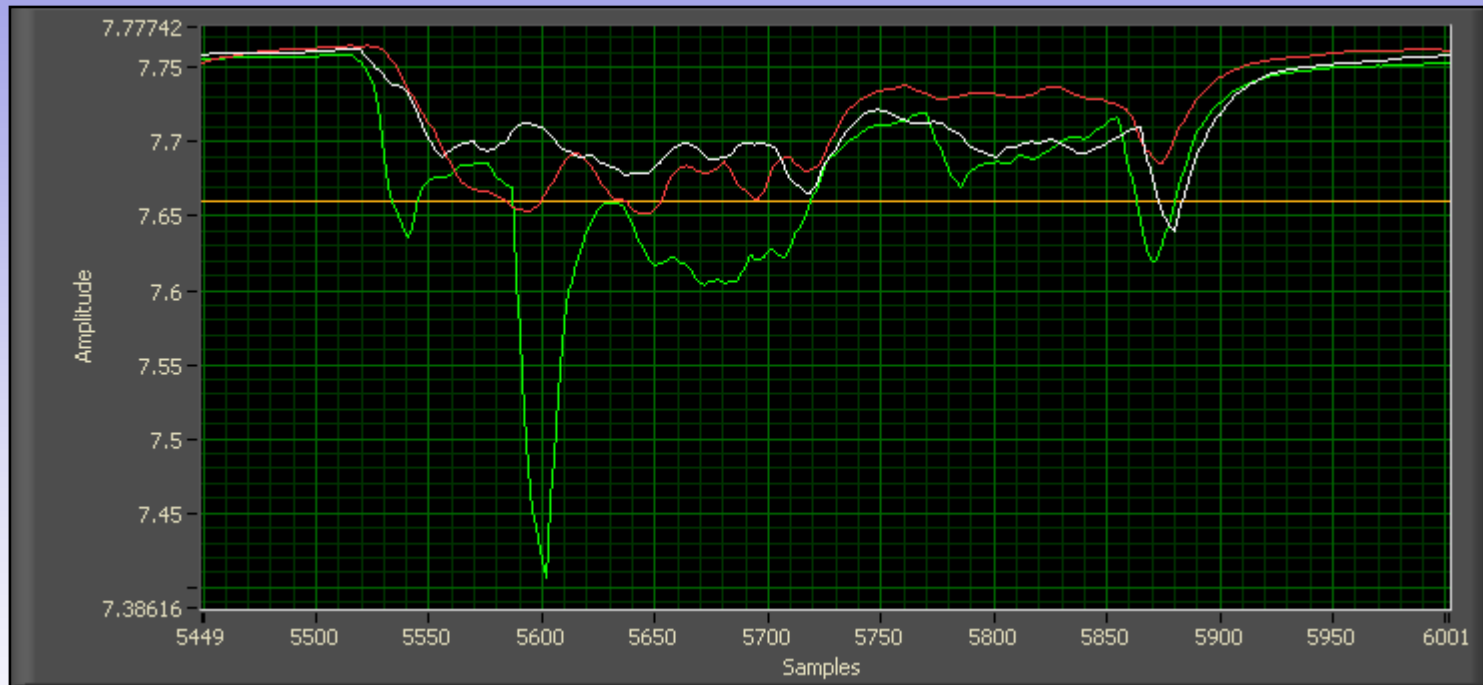


# FOCE Prototype Control GUI





# Example data – First FOCE Prototype Deployment – April 2005



White/red/green are top/mid/bottom sensors.  
Orange is the pH control set-point.



# Acknowledgements

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